# Fundamentals of Large-Eddy Simulation

### PALM group

#### Institute of Meteorology and Climatology, Leibniz Universität Hannover

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The Role of Turbulence ●○	The Reynolds Number 0	Classes of Turbulence Models	Concept of LES
The Role of Turbulence			

 Most flows in nature & technical applications are turbulent





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The Role of Turbulence			

- Most flows in nature & technical applications are turbulent
- Significance of Turbulence
  - Meteorology / Oceanography: Transport processes of momentum, heat, water vapor as well as other scalars
  - Health care: Air pollution
  - Aviation, Engineering: Wind impact on buildings, power output of windfarms





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#### Characteristics of turbulence

- non-periodical, 3D stochastic movements
- mixes air and its properties on scales between large-scale advection and molecular diffusion
- $\blacktriangleright$  non-linear  $\rightarrow$  energy is distributed smoothly with wavelength
- wide range of spatial and temporal scales



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Large eddies: 10<sup>3</sup> m (L), 1 h
 Small eddies: 10<sup>-3</sup> m (η), 0.1 s





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  - Large scales: shear and buoyant production
  - Small scales: viscous dissipation







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#### Energy-cascade

Large eddies are broken up by instabilities and their energy is handled down to smaller scales.





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The Reynolds Number			

## The Reynolds Number (Re)

$$rac{L}{\eta}pprox {\it Re^{3/4}}pprox 10^6$$
 (in the atmosphere)

 $\begin{array}{l} {\bf u} \mbox{ 3D wind vector} \\ \nu \mbox{ kinematic molecular viscosity} \\ L \mbox{ outer scale of turbulence} \\ U \mbox{ characteristic velocity scale} \\ \eta \mbox{ inner scale of turbulence} \\ (Kolmogorov \mbox{ dissipation length}) \end{array}$ 



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### The Reynolds Number (Re)

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$Re = rac{ \mathbf{u}\cdot abla\mathbf{u} }{  u abla^2\mathbf{u} } \hat{=} rac{LU}{ u}$	inertia forces viscous forces	$U$ change $\eta$ inn
		(Koln

**u** 3D wind vector  $\nu$  kinematic molecular viscosity L outer scale of turbulence U characteristic velocity scale  $\eta$  inner scale of turbulence (Kolmogorov dissipation length)





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# The Reynolds Number (Re)

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$$\Rightarrow$$
 Number of gridpoints for a 3D simulation:

$$\left(rac{L}{\eta}
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 (in the atmosphere)



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Classes of Turbulence Models			





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  - Limited computer resources (1996:  $\sim 10^8$ , today:  $\sim 10^{12}$  gridpoints, but  $\sim 10^{18}$  gridpoints needed, see prior slide).
  - 1 h simulation of 10<sup>9</sup> (2048<sup>3</sup>) gridpoints on 512 processors of the HLRN supercomputer needs 10 h CPU time.



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- Consequences:
  - DNS is restricted to moderately turbulent flows (low Reynolds-number flows).
  - Highly turbulent atmospheric turbulent flows cannot be simulated.



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Reynolds averaged (Navier-Stokes) simulation (RANS)





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  - Opposite strategy:
    - Applications that only require average statistics of the flow (i.e. the mean flow).
    - Integrate merely the ensemble-averaged equations.
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    - Parameterizations are very sensitive to large-eddy structure that depends on environmental conditions such as geometry and stratification → Parameterizations are not valid for a wide range of different flows.





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  - Consequence:
    - Not suitable for detailed turbulence studies.



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### Large eddy simulation (LES)

 Seeks to combine advantages and avoid disadvantages of DNS and RANS by treating large scales and small scales separately, based on Kolmogorov's (1941) similarity theory of turbulence.







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- Large eddies are explicitly resolved.
- The impact of small eddies on the large-scale flow is parameterized.
- Advantages:
  - Highly turbulent flows can be simulated.
  - Local homogeneity and isotropy at large *Re* (Kolmogorov's 1<sup>st</sup> hypothesis) leaves parameterizations uniformly valid for a wide range of different flows.



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#### Filtering

 Spectral cut at wavelength Δx.





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$$w = \overline{w} + w', \theta = \overline{\theta} + \theta'$$

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- The filter procedure removes the small scales from the model equations, but it produces new unknowns, mainly averages of fluctuation products.
  - ▶ eg. *w*′θ′





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- We do not have information about the variables (e.g., vertical wind component and potential temperature) on these small scales of their fluctuations.
- Therefore, these unknowns have to be parameterized using information from the resolved scales.
  - A typical example is the flux-gradient relationship, e.g.,

$$\overline{w'\theta'} = -\nu_{\rm h} \cdot \frac{\partial \overline{\theta}}{\partial z}$$

